

Base from U.S. Geological Survey, Port Moller, 1933 (revised 1988); Simeonof Island, 1963; Stebovak Bay, 1963 (revised 1981); Universal Transverse Mercator projection, zone 4

SCALE 1:250,000
CONTOUR INTERVAL 100 AND 200 FEET
DOTTED LINE REPRESENTS 100-FOOT CONTOURS
NATIONAL GEODETIC VERTICAL DATUM OF 1929

INDEX MAP SHOWING PORT MOLLER, STEBOVAK BAY, AND SIMEONOF ISLAND 1:250,000-SCALE QUADRANGLES AND STUDY AREA, ALASKA PENINSULA

Geologic mapped by L.M. Angeloni, 1983-85; J.E. Case, 1977-78, 1983-85; C.L. Connor, 1982; John Decker, 1985; R.L. Detterman, 1977-78, 1982-86; D.D. Dubois, 1985; E.M. Gamble, 1983-84; Stephanie Zurnicki, 1986; Louise Martynovich, Jr., 1982, 1984; J.W. Miller, 1982-85; M.A. Penobscot, 1983; Nora Shew, 1982; F.R. Weber, 1983-85; F.H. Wilson, 1977-78, 1982-86; M.E. Yount, 1977-78, 1982-85

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DESCRIPTION OF MAP UNITS

[Stratigraphic nomenclature and age assignment from Detterman and others, in press.]

SURFICIAL DEPOSITS AND SEDIMENTARY ROCKS

Qs Surficial deposits (Holocene and Pleistocene).—Unconsolidated, poorly to well-sorted, poorly to moderately well stratified sand, gravel, and silt. Also includes alluvial, colluvial, glacial, marine, lacustrine, and eolian deposits. Locally includes reworked volcanic-ash and debris-flow deposits.

Tmr Milky River Formation (Pliocene).—Volcanogenic nonmarine sedimentary rocks and interbedded flows and sills. Upper part of unit contains numerous porphyritic andesitic lava flows, lahars, and tuff beds interbedded with sedimentary rocks. Lower part consists almost entirely of coarse-grained, highly crossbedded and channelled, fluvial volcanic sandstone and cobble-boulder conglomerate that has clasts composed mainly of volcanic debris.

Tbl Bear Lake Formation (late Miocene).—Inner neritic marine and nonmarine (Wieshart, 1971; Nilsen, 1984) sandstone, conglomerate, siltstone, and shale. Sandstone is moderately well sorted and grains are moderately well rounded. Conglomerate horizons are made up of well-sorted clasts, of which about 40 to 55 percent are quartz and chert. 20 to 30 percent are volcanic fragments, 10 to 15 percent are granitic clasts, and remainder is lithic sedimentary clasts. Unit is abundantly fossiliferous.

Tu Unga Formation (middle Miocene to late Oligocene).—Volcaniclastic sedimentary and volcanic rocks overlying coal-bearing siltstone and shale. Volcanic rocks are dominant in upper parts where carbonaceous shale and coal are confined to lower part. Sandstone and conglomerate are composed of poorly sorted volcanic debris and are poorly consolidated. Many conglomerate beds are lahar deposits.

Tbe Belkofski Formation (middle and early Miocene).—Mainly tuffaceous, volcaniclastic sandstone, siltstone, and conglomerate and interbedded breccia (McLennan, 1979). Rocks are dominantly red, pink, and purple and very well indurated. Potassium-argon (K-Ar) dates from the Belkofski Formation on Dagle Island is 11.780±0.41 Ma (from Shew, unpublished, 1988). If mapping is correct, then some of the Belkofski is of Miocene age.

Ts Stebovak Formation (early Oligocene and late Eocene).—Upper part is olive-gray and yellowish-brown sandstone rich in unlaminated volcanic debris and was deposited in a nearshore, shallow-water, shelf environment. Lower part is dark brown laminated siltstone and shale deposited as a deep-water turbidite and commonly shows graded bedding and rip-up clasts in sandstone interbeds. Mudstone distributed throughout upper part is characteristic of water depths no greater than 30 to 50 m (Louie Martynovich, Jr., written commun., 1983-1986).

Tt Tolstoi Formation (middle Eocene to late Paleocene).—Pale-yellowish-brown to gray-green interbedded sandstone, conglomerate, siltstone, and shale. Lithic clasts in conglomerates are dominantly granitic detritus and arkosic detritus that contains 20 to 30 percent volcanic clasts. Most volcanic clasts are not fresh appearing and, in association with the granitic and arkosic detritus, suggest a Mesozoic source rather than contemporaneous magmatic activity, in sharp contrast to most overlying units (Detterman and others, in press). In type area just east of Pavlov Bay, characteristic lithologies indicate shallow-marine sedimentation, which is succeeded northward by delta plain and fluvial deposits, mainly of braided-stream type, that are typical for the major part of Tolstoi Formation.

Kc Chignik Formation (late Cretaceous).—Dominantly a light-olive-gray to olive-gray sandstone and interbedded olive-gray to olive-black siltstone and conglomerate of multicolored chert, white quartz, granitic rocks, and minor volcanic rocks. Unit is cyclic nearshore marine. Ideal flat, nonmarine flood plain, and fluvial deposits (Fairchild, 1977; Detterman, 1978). Nonmarine parts may contain coal beds as much as 2 m thick. Marine fossils, mainly bryozoans, indicate a late Campanian to early Maestrichtian age (J.W. Miller, written commun., 1983-1986).

Kh Hoodoo Formation (late Cretaceous).—Typically dark-gray to black, highly indurated and bedded, splitting to pencil fracturing siltstone and fine-grained sandstone. Becomes more sandy upwards, and in some areas contains channel conglomerate that has clasts of pholitic and volcanic rocks, chert, and quartz. Sandstone beds range from 0.3 to 1 m thick and siltstone and shale beds range from 1 to 2 m thick and have individual layers as thin as 1 cm (Detterman and others, 1981). Sparse megafossils are common in age of late Campanian to early Maestrichtian (J.W. Miller, written commun., 1983-85). Depositional environment is characteristic of lower to upper slope of a submarine fan; structures imply submarine slumping and turbidity current flow.

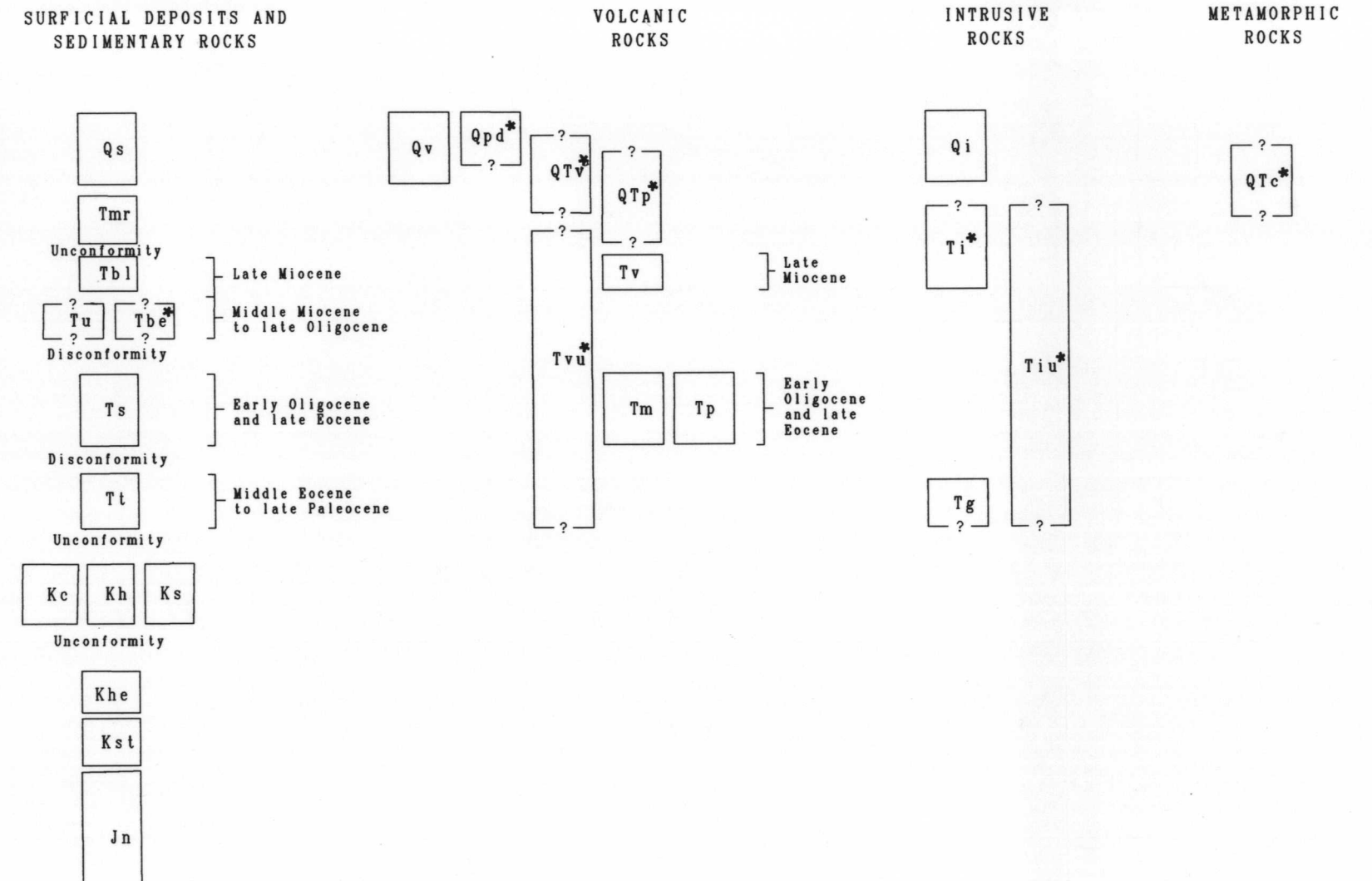
Xs Shumagin Formation (late Cretaceous).—Interbedded graywacke, siltstone, mudstone, and shale. Sandstone is medium to very fine grained, highly indurated and medium light gray to medium dark gray (Detterman, 1978). Siltstone is light graywacke, fine grained, and contains abundant shale and siltstone chips (Moore, 1974; Detterman and others, 1981). Sparse megafossils are common in age of late Campanian to early Maestrichtian (J.W. Miller, written commun., 1983-85). Depositional environment is characteristic of lower to upper slope of a submarine fan; structures imply submarine slumping and turbidity current flow.

Herendeen Formation (early Cretaceous).—Primarily uniform fine-grained, medium grained, and dark gray to pale yellowish-brown to brown fresh broken surfaces. Facies fragment from a thin bedded, medium grained, and dark gray to pale yellowish-brown to brown fresh broken surfaces. Facies fragment from a thin bedded, medium grained, and dark gray to pale yellowish-brown to brown fresh broken surfaces. Facies fragment from a thin bedded, medium grained, and dark gray to pale yellowish-brown to brown fresh broken surfaces.

Stanukovich Formation (early Cretaceous).—Light-olive gray siltstone containing thin-bedded sandstone and siltstone. Unit is typically olive gray siltstone that has numerous calcareous nodules and concretions. Upper unit erodes readily, is typically well exposed, and contains few age-diagnostic fossils. Lower unit has an abundant megafossil, particularly the bryozoan, which indicates a Paleocene age (Detterman and others, 1981; J.W. Miller, written commun., 1982-86).

Naknek Formation (late Jurassic).—Dominantly medium-gray, fine- to medium-grained arkosic sandstone and siltstone. Fresh biotite and hornblende are a minor but important component of the sandstone. Dark-gray siltstone of the Sing Harbor Siltstone Member (Detterman and others, in press) and light-gray arkosic sandstone containing magnetite laminae and thin beds of conglomerate of the Northeast Creek Sandstone Member (Detterman and others, in press) are present in Port Moller area and northwest of Supai Lake. Abundant megafossil collections from formation indicate an age of Ordovician to Triassic (Detterman and Reed, 1984; p. 133; J.W. Miller, written commun., 1982-88).

CORRELATION OF MAP UNITS



INTRUSIVE ROCKS

QI Younger intrusive rocks, undivided (Holocene and Pleistocene).—Hypabyssal dike plugs and dikes of Quaternary volcanic centers, particularly Trader Mountain and Mount Dana. Porphyritic dike at Trader Mountain has a K-Ar age of 0.865±0.05 Ma (P.H. Wilson and Nora Shew, unpub. data, 1988).

TI Older intrusive rocks, undivided (Pliocene and late Miocene).—Medium- to coarse-grained, equigranular granodiorite to quartz diorite plutons and stocks containing hornblende, biotite, and pyroxene. Typically display well-developed hornfels zones and sporadic hypothermal alteration in surrounding country rocks. Potassium-argon ages range from 9.43±0.26 to 8.21±0.14 Ma (P.H. Wilson and Nora Shew, unpub. data, 1990).

Thu Intrusive rocks, undivided (Tertiary).—Small intrusive bodies of quartz diorite or diorite, which are typically hypabyssal and contain phenocrysts of pyroxene or hornblende in a fine-grained groundmass.

Tg Granitic rocks (Pliocene).—Medium-grained biotite granodiorite and quartz monzonite plutons that have hypidionomorphic granular textures and local development of potassium feldspar phenocrysts as much as 1 cm long. Crop out on Nagai, Big and Little Konliji, Simeonof, Chernabura and Bird Islands. Potassium-argon ages (Burk 1965; Moore, 1974; Kemle and Turner, 1976) range from 65.6±3.3 to 57.9±1.8 Ma (recalculated using constants of Steiger and Jager, 1977).

METAMORPHIC ROCKS

QTc Contact-metamorphosed rocks (early Quaternary or late Tertiary).—Contact metamorphosed and hydrothermally altered rocks in mountains east of Mount Dana. Rocks probably consist of the Hoodoo Formation (Kh), undivided Tertiary volcanic rocks (Ti), and Chignik Formation (Kc). Unit is well indurated, very fine grained, and intensely fractured; sulfide mineralization is common, and resultant iron staining is ubiquitous on weathered surfaces.

Legend:
Contact—Dashed where approximately located; dotted where concealed.
Fault—Dashed where approximately located; dotted where concealed.
Thrust or high-angle reverse fault—Dashed where approximately located; dotted where concealed. Swath on upper plate.
Volcanic crater

INTRODUCTION

Geologic mapping of the Port Moller, Stebovak Bay, and Simeonof Island 1:250,000-scale quadrangles on the Alaska Peninsula was conducted as part of the Alaska Peninsula Geologic Map (APGMAP) project, a cooperative effort of the U.S. Geological Survey and the Alaska Department of Natural Resources. The project was initiated in 1977 by the U.S. Geological Survey (USGS) and the Alaska Department of Natural Resources (ADNR). The project was completed in 1988. The project was a cooperative effort of the U.S. Geological Survey and the Alaska Department of Natural Resources. The project was initiated in 1977 by the U.S. Geological Survey (USGS) and the Alaska Department of Natural Resources (ADNR). The project was completed in 1988. The project was a cooperative effort of the U.S. Geological Survey and the Alaska Department of Natural Resources.

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GENERALIZED GEOLOGIC MAP OF THE PORT MOLLER, STEBOVAK BAY, AND SIMEONOF ISLAND QUADRANGLES, ALASKA PENINSULA, ALASKA

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